TRACING THE ECOPHYSIOLOGY OF UNGULATES AND PREDATOR-PREY RELATIONSHIPS IN AN EARLY PLEISTOCENE LARGE MAMMAL COMMUNITY

Paul Palmqvist a,*, Juan A. Pérez-Claros a, Christine M. Janis b, Darren R. Gröcke c

a Departamento de Ecología y Geología, Facultad de Ciencias, Campus Universitario de Teatinos, 29071-Málaga, Spain

b Department of Ecology and Evolutionary Biology, Brown University, Providence, RI 02912

c Department of Earth Sciences, Durham University, Science Laboratories, South Road, Durham DH1 3LE, UK

* Corresponding author. Departamento de Ecología y Geología, Facultad de Ciencias. Campus de Teatinos s/n, E-29071-Málaga. E-mail address: ppb@uma.es (P. Palmqvist), Fax: (+34) 952137386

Abstract

Research into the reconstruction of ancient communities in terms of dietary regimes, habitat preferences and ecological interactions of species has focused predominantly on biogeochemistry or ecomorphology alone and not in combination. The Venta Micena site (Orce, Guadix-Baza basin, SE Spain) has an early Pleistocene vertebrate assemblage with exceptional biomolecular preservation. Collagen was successfully extracted from 77 bone samples of 18 species of large mammals, which allowed analyses of carbon- and nitrogen-isotopes. δ13C, δ15N and δ18O ratios combined with ecomorphological indexes provide interesting clues on the autecology and palaeophysiology of extinct species, which help in deciphering aspects of community trophic structure and predator-prey interactions. Specifically, morphometric ratios (e.g., hypsodonty index and relative length of the lower premolar tooth row; Palmqvist et al., 2003) allow classifying the ungulates among grazers
(Equus altidens, Bison sp., Praeovibos sp., Hemitragus albus, Hippopotamus antiquus, and Mammuthus meridionalis), mized-feeders (Soergelia minor and Pseudodama sp.) and browsers (Stephanorhinus sp. and Praemegaceros cf. verticornis). However, δ¹³C values reveal that these ungulates consumed exclusively C₃ plants and significant differences in isotopic values between perissodactyls (monogastric, hindgut fermenters) and ruminants (foregut fermenters) must reflect physiological differences related to their rates of methane production and digestive efficiency. δ¹⁸O ratios allow the interpretation of the dietary water source of these species, suggesting that fallow deer Pseudodama sp., goat H. albus and ovibovine S. minor obtained a significant fraction of their metabolic water from the vegetation consumed. Carnivore species have higher δ¹⁵N values than herbivores, which records the isotopic enrichment expected with an increase in trophic level. However, the unexpectedly high δ¹⁵N values of hippo H. antiquus and muskoxen Praeovibos sp. suggest that these ungulates predominantly consumed aquatic plants and lichens, respectively.

Inferences on predator-prey relationships within this ancient community, derived from the dual linear mixing model, indicate resource partitioning among sympatric predators, suggesting that sabre-tooth Megantereon whitei and jaguar Panthera cf. gombaszoegensis were ambushers of forest environments while sabre-tooth Homotherium latidens and wild dog Lycaon lycaonoides were coursing predators in open habitat. The giant, short-faced hyena Pachycrocuta brevirostris scavenged the prey of these hypercarnivores.

Keywords: Mammals; Ecomorphology; Biogeochemistry; Pleistocene; Venta Micena; Orce

1. Introduction: the early Pleistocene locality of Venta Micena

Venta Micena lies near the village of Orce (Granada, SE Spain) in the eastern sector of the Guadix-Baza Basin (37°44′15″N, 2°24′9″W, elevation 974.5 m; Fig. 1). This sedimentary basin was characterized by interior drainage from the end of the Miocene to middle-late Pleistocene times, which facilitated the preservation of Plio-Quaternary large mammal assemblages in swamppy
and lacustrine sediments. The site is dated by biostratigraphy to the lower Pleistocene (Arribas and Palmqvist, 1999), with an age estimated in ~1.5 Ma.

The Venta Micena stratigraphic column shows alternate micrite limestone, calcilutitic, lutitic, silty, and marly levels (Fig. 1). The main excavation quarry (VM-2 level, Quarry 3 and drillings 1-4, ~320 m²; Palmqvist and Arribas, 2001) is located within the upper part of the section, in an 80-120 cm thick limestone stratum undisturbed by tectonic activity, composed of homogeneous and porous micrite sediments (98-99% CaCO₃) that can be followed across ~2.5 km in the Orce area (Arribas and Palmqvist, 1998). The lower half of the stratum has carbonate nodules (5-20 cm thick), mud banks and fossil shells of eurythermal freshwater molluscs. During this lacustrine stage, micrite was precipitated in a shallow (<10 m), well-oxygenated water sheet not subject to eutrophic conditions, as indicated by the absence of pyrite and carbonate facies rich in organic matter. Above this level there is a 4-15 mm thick calcrete palaeosol, which records a major retreat of the Pleistocene lake and represents a swampy biotope with wide emerged zones (~4 km width) and shallow ponds (<1 m depth, 2-20 m diameter). The upper half of the stratum is composed of micrite sediments showing root marks and mud cracks at the bottom, which record the rise of the lake level, and preserves a high density of fossil bones.

The large mammals assemblage is composed of ~5,800 identifiable skeletal remains from 225 individuals belonging to 21 taxa of large (>5 kg) mammals and ~10,000 unidentifiable bone shafts and cranial fragments. Herbivorous taxa dominate the assemblage in number of identifiable specimens (NISP) and estimates of minimal number of individuals (MNI). The surface of the skeletal remains is not abraded and the longitudinal axes of long bones show no preferred orientation, which indicates that they were not transported by fluvial processes prior to deposition. Furthermore, the ratio of isolated teeth to vertebrae (0.94:1) is close to the value expected in the absence of hydrodynamic sorting (1:1) and the frequencies of bones grouped according to their potential for dispersal by water (i.e., Voorhies’s groups) are similar to those in the mammalian skeleton (Arribas and Palmqvist, 1998). Analysis of weathering stages indicates a short time of
exposure before burial, less than one year in most cases. Analysis of mortality patterns deduced for ungulate species from juvenile/adult proportions reveals that most skeletal remains were scavenged by the giant, short-faced hyena *Pachycrocuta brevirostris* from carcasses of animals hunted selectively by hypercarnivores (Palmqvist et al., 1996). Taphonomic analysis shows that the hyenas transported ungulate carcasses and body parts to their maternity dens as a function of the mass of the ungulates scavenged. The fracturing of major limb bones in the dens was also highly selective, correlating well with their marrow contents and mineral densities (Palmqvist and Arribas, 2001).

Palaeoecological analyses include inferences on the life style and preferred habitat of extinct taxa (palaeoautecology) and the reconstruction of past ecological associations (palaeosyneceology) (Damuth, 1992). Once the preservational completeness of the fossil assemblage has been evaluated with taphonomic analysis, it is necessary to infer the autecological characteristics of the species prior to synecological analysis at the community level. Autecological properties of extinct taxa may be reconstructed through: 1) ecomorphological inferences on functional adaptations for feeding behaviour and types of locomotion; 2) biogeochemical analyses for reconstructing dietary niches, habitat preferences and palaeotemperatures; and 3) studies of the sedimentary context and taphonomic attributes of the fossils as well as on their distribution across facies (Wing et al., 1992; Palmqvist et al., 2003; Soligo and Andrews, 2005). Once this goal is achieved, species can be distributed among size classes and ecological categories, and the relative frequencies of these categories in the assemblage are compared with those in modern ecosystems (Andrews et al., 1979; Reed, 1998; Mendoza et al., 2005).

Feeding preferences of extinct mammals can be addressed using biogeochemical markers such as stable-isotopes and trace-elements, as well as from the comparative study of their craniodental morphology, because several features of the skull, mandible and dentition are indicative of diet (see reviews in MacFadden, 2000; Williams and Kay, 2001; Mendoza et al., 2002; Palmqvist et al., 2003). In herbivores, the hypsodonty index (*HI*, unworn molar crown height divided by molar width) discriminates between grazers (>75% grass in diet), which feed upon
grasses with high silicophytolith contents and have hypsodont, high-crowned molars, from browsers
(<25% grass in diet), which consume succulent leaves and have brachydont, low-crowned teeth.

However, although HI is probably the best single variable correlated with diet in living ungulates
(and thus of the best use for predicting the diet of extinct ones), in some species molar crown height
alone may be insufficient to determine their feeding preferences (Fortelius and Solounias, 2000;
Mendoza et al., 2002). Muzzle shape also provides information on diet, as it reflects the adaptations
related to the “cropping mechanism”, including the shape of the premaxilla and the relative
proportions of the incisor teeth (Janis and Ehrhart, 1988; Solounias and Moelleken, 1993; Pérez-
Barbería and Gordon, 2001): browsers have narrow muzzles consisting of a rounded incisor arcade
with the first incisor generally larger than the third, while grazers have broad muzzles with
transversely straight incisor arcades, showing equal or sub-equal sized teeth. There are, however,
some second-order differences related to the phylogenetic legacy: equids have relatively narrower
muzzles than grazing ruminants of similar body size. In addition, different ungulate groups have
adopted different solutions when faced with the same ecological specialization: for example, the
lower premolars are enlarged in grazing perissodactyls, but grazing ruminants and camelids show
the opposite trend. This difference is probably due to differences in the way food is orally processed
in foregut and hindgut fermenters (Mendoza et al., 2002).

In carnivores, craniodental features related to diet include the morphology of the upper
canine, the size of the trigonid blade and the talonid basin in the lower carnassial, and the shape of
the glenoid and angular processes in the mandible, which reflect the moment arms for jaw adductor
muscles (Van Valkenburgh, 1988; Biknevicius and Van Valkenburgh, 1996). These variables help
in discriminating among hypercarnivores, bone-crackers and omnivores. Relevant features of the
postcranial skeleton include the brachial and crural indexes (i.e., radius length divided by humerus
length and tibia length divided by femur length, respectively), the ratio of phalanx length to
metacarpal length, the biceps brachii leverage index, and cross-sectional geometric properties of
major limb bones (Van Valkenburgh, 1985; Anyonge, 1996; Lewis, 1997). These variables estimate
different aspects related to habitat preferences and hunting techniques: for example, the brachial and crural indexes are useful for discriminating between predators that ambush their prey in forested environments and those that pursue it in open habitat (Palmqvist et al., 2003).

2. Sedimentary geochemistry and palaeoenvironmental inferences

The trace-element and stable-isotope chemistry of lake waters is a sensitive monitor of climate in arid and semi-arid regions (see review in Hu et al., 1998). Data of trace-element abundance and stable-isotope ratios in the stratigraphic section of Venta Micena (Table 1; Figs. 1A-B) show several key relationships. For example, sodium concentrations decrease systematically from the base of the section, reach their lowest values in the middle of the section, and then suddenly rise to high values below the palaeosoil (Fig. 1A). If sodium concentrations are assumed as representative of lake salinity levels, they show that the middle part of the section witnessed an increase in the water level and/or a reduction of salinity, although the salinity level dramatically increased just prior to palaeosoil development, evidencing the lowering of the water table. Iron and manganese concentrations decrease from level VM-1 to level VM-2, suggesting a shift from a restricted, stratified water column (decreased oxygen levels, stagnation) to a more open, well-oxygenated water column. This evidence supports the increase in lake-level in the middle of the section, as indicated by sodium concentrations, which resulted in increased water supply and, thus, circulation.

Magnesium and strontium concentrations decrease and increase, respectively, through the whole stratigraphic section (Fig. 1A). Magnesium and strontium concentrations are under saturated with respect to the minerals commonly precipitated within lakes and for this reason both elements have been used for reconstructing changes in water lake salinity levels (Chivas et al., 1985). Magnesium also correlates negatively with water temperature (Chivas et al., 1986). However, on the basis of these bulk-rock analyses, a similar assumption would be difficult here, because magnesium and strontium are negatively correlated in the Venta Micena section.
Oxygen-isotope analyses of the bulk-rock samples show a general decrease to more negative δ¹⁸O values through the section (Fig. 1B), paralleling the magnesium concentration record. This indicates that magnesium and δ¹⁸O are negatively correlated with the palaeotemperature of the lake waters and, thus, an overall warming from level VM-1 to the development of the palaeosoil at level VM-2. Sodium and strontium concentrations should be negatively correlated if strontium measures palaeosalinity levels, what is not reflected in the Venta Micena section. However, in a palaeohydrochemical study of a nearby early Pleistocene shallow lacustrine section from Orce, Anadón and Julià (1990) found lower Sr/Ca values in ostracod shells from sands deposited during saline water phases than in those from the overlying carbonate sequences formed under lower salinity conditions; such unexpected values were interpreted as the result of major changes in the chemical composition of the water in shallow swamped areas of a hydrologically complex lake. A subsequent study (Anadón et al., 1994) revealed higher δ¹⁸O ratios in ostracod shells from intervals with a saline fauna than in those with a freshwater fauna, what is also recorded in the bulk-rock analyses of the Venta Micena section (Fig. 1B). According to Anadón et al. (1994), this would correspond to an alternation of concentration/dilution phases in a shallow lacustrine sequence that correlates with the climatic cycles described in synchronous ocean basin records from the late Matuyama chron. Anadón et al. (1994) also found a covariant trend in δ¹³C and δ¹⁸O values from ostracod calcite, which indicates that the ostracods lived in a closed lacustrine system. δ¹³C and δ¹⁸O ratios are also correlated in the Venta Micena section (Fig. 1B).

Bulk-rock δ¹³C ratios are, however, an archive of more difficult interpretation. They show the maximum value in the marly level at the base of the section, fluctuate in the silty, lutitic and calcilutitic levels placed in the section from 0.85 m to 2.5 m, and then show a slight decrease in the micrite levels (Fig. 1B). Organic residues in modern soils reflect the δ¹³C of the overlying flora (Koch, 1998). Because its carbon is derived from soil CO₂, the δ¹³C of soil carbonate is strongly correlated to that of soil organic matter. Atmospheric CO₂ has a higher δ¹³C value (−6.5‰) than both C₃ plants (−26‰) and C₄ plants (−12‰), contributing to soil CO₂ near the surface; however,
the CO$_2$ >30 cm deep in soils with moderate to high respiration rates is largely supplied by plant
decay and root respiration. Both processes generate CO$_2$ isotopically similar to organic matter.

Diffusion of CO$_2$ from the soil to the atmosphere leads to a $\delta^{13}$C enrichment of +4.5‰ for CO$_2$ at
depth in a soil relative to soil organic matter. Finally, temperature-dependent fractionation
associated with precipitation of calcite sum to a $\delta^{13}$C increase of +10.5‰ (Koch, 1998). As a
consequence, modern carbonates forming below 30 cm depth have $\delta^{13}$C values ~15‰ higher on
average than those of organic matter: −11‰ for soils with C$_3$ overlying flora and +3‰ for soils in
which organic matter is supplied by C$_4$ plants (Koch, 1998). The range of $\delta^{13}$C ratios measured in
the Venta Micena stratigraphic section (−4.1‰ to −7.4‰) lies between both values, suggesting a
mixed vegetation of C$_3$ and C$_4$ plants. However, $\delta^{13}$C values for bone collagen of grazing ungulates
(see below) show the absence of C$_4$ grasses in their diet. This indicates that other factors apart from
changes in primary productivity and respiration in the water column may also have been involved in
determining bulk-rock $\delta^{13}$C ratios; for example, under higher pressures of atmospheric CO$_2$, more
of the CO$_2$ at depth in soils would be derived from the atmosphere, increasing the difference in $\delta^{13}$C
values between soil carbonate and organic matter (Koch, 1998).

In a recent study of a 356-m-thick composite section of the Guadix-Baza basin that ranges
from the late Pliocene to the middle Pleistocene, Ortiz et al. (2006) interpreted the $\delta^{13}$C and $\delta^{18}$O
profiles as reflecting changes in temperature, the evaporation/infill ratio in the water bodies and the
amount of rain. Specifically, they concluded that high $\delta^{13}$C and $\delta^{18}$O values were associated with
warm and dry regimes, whereas low $\delta^{13}$C and $\delta^{18}$O values correlated with cold and humid episodes,
which caused more vegetation biomass and, therefore, an increase in the input of isotopically light
carbon.

Strontium isotopes ($^{87}$Sr/$^{86}$Sr) can be used for deriving the palaeosalinity record of ancient
environments if independent constraints on the system’s hydrologic parameters (i.e., evaporation,
precipitation, fluvial and ocean exchange fluxes) are available (e.g., salinity estimates provided by
lithology and faunal assemblages; Flecker et al., 2002). $^{87}$Sr/$^{86}$Sr ratios of river waters are similar to
those of terrestrial plants and there are no significant differences in $^{87}$Sr contents between grasses and trees (Hoppe et al., 1999). In herbivores, the $\delta^{87}$Sr value of bioapatite equals the average ratio of the vegetation ingested, which in turn monitors the soluble strontium in soils, derived from weathering and precipitation. Environmental $^{87}$Sr/$^{86}$Sr ratios vary with differences in atmospheric input as well as with differences in bedrock age and composition (Price et al., 1985; Miller et al., 1993). Due to this reason, variations in $\delta^{87}$Sr have been used for reconstructing migratory behaviour in a variety of vertebrates, including proboscideans (Koch et al., 1995; Hoppe et al., 1999; Hoppe, 2004). Concerning the stratigraphy of Venta Micena (Fig. 1B), bulk-rock $^{87}$Sr/$^{86}$Sr ratios are relatively uniform in the lower part of the section, with the only exception of a decrease in level VM-1. This reflects deposition under conditions of hydrological stability. The upper carbonate samples, however, show a significant decrease in $^{87}$Sr/$^{86}$Sr proportions, which reflects an increase in river or groundwater input that translated in the rising of the lake’s table in the lacustrine levels.

3. Stable-isotope analyses of the Venta Micena fauna

Stable isotopes have proved useful in determining the dietary niches of extinct mammals, providing detailed ecological and environmental reconstructions. Published carbon-, nitrogen- and oxygen-isotopes of collagen and hydroxylapatite from 18 species of large mammals identified in the Venta Micena assemblage (Gröcke et al., 2002; Palmqvist et al., 2003; N = 65) and results obtained from additional samples (N = 50; Table 2) of juvenile individuals and species not sampled in the previous study (e.g., Praeovibos sp., Hystrix major, Panthera cf. gombaszoegensis, and Ursus etruscus) are analyzed here to determine their dietary niches and predator-prey relationships. Collagen was successfully extracted from 77 bone samples, which allowed analyses of carbon- and nitrogen-isotopes. Oxygen-isotopes were retrieved from 115 bone and tooth hydroxylapatite samples. The precision for stable-isotope analysis was 0.1‰ for both carbon and oxygen, and 0.2‰ for nitrogen. Carbonate carbon and oxygen isotopic analyses were performed using a Fison Optima isotope-ratio mass-spectrometer, with a common acid bath system in the Stable-Isotope Biogeochemistry Laboratory at McMaster University. Samples were reacted with 100% phosphoric
acid at 90ºC. Collagen carbon and nitrogen isotopic analyses were performed at the Stable-Isotope Biogeochemistry Laboratory at McMaster University using a Thermo-Finnigan DeltaPlus XP coupled with a Costech elemental analyzer.

Stable isotopes are useful palaeobiological tracers because, as a result of mass differences, different isotopes of an element have different thermodynamic and kinetic properties, leading to measurable isotopic partitioning during physical and chemical processes, which labels the substances with distinct isotopic ratios (Gröcke, 1997a; Koch, 1998). Isotopic ratios are reported as parts per thousand (‰) of deviation from a standard, using the δ notation, where:

\[
\delta X = \left(\frac{R_{\text{sample}}}{R_{\text{standard}}} - 1\right) \cdot 1000,
\]

where \(X = C, N\) or \(O\), and \(R = ^{13}C/^ {12}C, ^{15}N/^ {14}N\) or \(^{18}O/^ {16}O\). \(R_{\text{sample}}\) and \(R_{\text{standard}}\) are the high-mass to low-mass isotope ratios of the sample and the standard, respectively. Common standards for \(\delta^{13}C\), \(\delta^{15}N\) and \(\delta^{18}O\) are Peedee belemnite (PDB), atmospheric \(N_2\) and standard mean ocean water (SMOW), respectively.

The measurement of carbon- and nitrogen-isotope ratios in an animal's bone collagen provides an indication of aspects of its overall diet for the last few years of life (De Niro and Epstein, 1978). Apart from a report on bone collagen preserved in Late Cretaceous dinosaurs (Ostrom et al., 1993), original carbon- and nitrogen-isotope compositions have been retrieved from organic residues in fossils as old as 200 ka (Jones et al., 2001), though adequate preservation in such ancient specimens is rare (Bocherens et al., 1996a; Gröcke, 1997a). Thus, the extraction of collagen from 77 out of 105 fossil bone samples of Venta Micena, a locality with an age of ~1.5 Ma, constitutes an example of unusual biomolecular preservation. It is worth mentioning that other fossil proteins (e.g., albumin and immunoglobin) have also been detected in fossil samples from this site using immunological techniques (Torres et al., 2002).

Several methods are available to identify alteration of collagen, including analysis of C:N ratios (between 2.9 and 3.6) and amino acid composition (Gröcke, 1997a; Richards et al., 2000;
Drucker et al., 2003). These criteria allow for the identification and exclusion of collagen that is heavily degraded and/or contaminated. This is not the case at Venta Micena, because C:N ratios of the collagen material extracted averaged 3.18 (Table 2) and the amino acid composition from four specimens is similar to that of bone collagen in modern mammals, indicating good preservation (Palmqvist et al., 2003; Fig. 6).

3.1. Carbon-isotopes and palaeodiet

Terrestrial plants can be divided into two main groups on the basis of their photosynthetic pathway (Edwards and Walker, 1983): 1) C₃ plants, which follow the Calvin-Benson cycle (atmospheric CO₂ is fixed through the reductive pentose phosphate pathway); and 2) C₄ plants, which use the Hatch-Slack cycle (C₄-dicarboxylic acid pathway). C₃ plants are all trees and bushes, temperate shrubs and grasses adapted to cool/moist climate and/or high altitude, whereas C₄ plants include tropical, arid-adapted grasses. All plants take up ¹²CO₂ in preference to ¹³CO₂, but are important differences in their isotopic composition related to their carboxylating enzymes (Smith and Epstein, 1971; Gröcke, 1997b; Koch, 1998). C₃ plants use the ribulose carboxylase and have an average δ¹³C value of −26.0 ± 2.3‰ (range: from −35‰ in closed canopy to −20‰ in open areas exposed to water stress). C₄ plants use the phosphoenolpyruvate carboxylase, which discriminates less effectively against ¹³CO₂, showing a mean δ¹³C value of −12.0 ± 1.1‰ (range: −19‰ to −8‰).

When plants are consumed by herbivores, the plant carbon is incorporated into their skeletal tissues with some additional fractionation. The difference between the δ¹³C value of the animal's diet and that subsequently incorporated into bone collagen (δ¹³Cₑ) translates in an average increase per trophic level of +1.0‰ (range: 0‰ to +4.5‰) and a similar enrichment is recorded in carnivores (see review in Bocherens and Drucker, 2003). The isotope enrichment factor for biogenic apatite is higher, +14.1‰ on average for ungulates (Cerling et al., 2003).
3.2. Nitrogen-isotopes and trophic level

The nitrogen-isotope composition of collagen in mammals records their position in the trophic web, since each trophic level above herbivore is indicated by a mean increase in $\delta^{15}$N of ~3.4‰ (range: 1.7‰ to 6.9‰; Robinson, 2001; Bocherens and Drucker, 2003; Vanderklift and Ponsard, 2003). Soil synthesis of nitrogen, the diet of the animal (i.e., if it consumes N$_2$-fixing or non-N$_2$-fixing plants) and nitrogen metabolism are the primary factors that affect the $\delta^{15}$N value expressed in herbivore collagen (Sealy et al., 1987; Virginia et al., 1989; Bocherens et al., 1996b; Gröcke et al., 1997; Koch, 1998; Robinson, 2001). Herbivores from closed environment show lower $\delta^{15}$N values than those from open grassland because of soil acidity in dense forest. Plants that fix nitrogen have $\delta^{15}$N values that cluster close to the atmospheric N$_2$ value of 0‰, whereas those that do not fix nitrogen and use other sources (e.g., soil NH$_4^+$ and NO$_3^-$) show a wider range of values. Therefore, animals consuming N$_2$-fixing plants will generally exhibit $\delta^{15}$N values between 0‰ and 4‰, while herbivores feeding on non-N$_2$-fixing plants will show $\delta^{15}$N values comprised between 2‰ and 8‰. Plants near marine or salt-affected areas show enrichment in $\delta^{15}$N values and deep-rooted plants are enriched over those with shallow roots.

The effects of nitrogen metabolism in mammals are very important. Higher $\delta^{15}$N values are observed in mammals inhabiting arid regions, due to ecophysiological differences in nitrogen metabolism associated with adaptations for drought tolerance: under conditions of increased aridity, mammals concentrate urine and excrete concentrated urea, subsequently causing elevated $\delta^{15}$N values (Gröcke, 1997a; Koch, 1998; Schwarcz et al., 1999). In general, higher $\delta^{15}$N values are found in perissodactyls (monogastric, hindgut-fermenting herbivores) than in foregut ruminants (Gröcke and Bocherens, 1996), but the cause behind these elevated values is not clearly understood. Ruminants have a distinct process of nitrogen cycling, where some waste urea is dumped into the rumen, and they are thus less water dependent than monogastric herbivores. Elevated $\delta^{15}$N levels
may also be indicative of young suckling animals, due to the ingestion of nutrient-enriched milk (Gröcke, 1997a; Jenkins et al., 2001).

### 3.3. Oxygen-isotopes and water requirements

Oxygen-isotope values in enamel and bone apatite reflect prevailing climatic conditions (e.g., palaeotemperature; Koch et al., 1989; Ayliffe et al., 1992), but they also allow the interpretation of the dietary water source of a local fauna (Sponheimer and Lee-Thorp, 2001; Harris and Cerling, 2002). The oxygen-isotope composition of apatite is a function of three main oxygen sources: atmospheric O$_2$, liquid water and oxygen bound in food (Bryant and Froelich, 1995; Bryant et al., 1996; Kohn, 1996; Kohn et al., 1996). Unlike atmospheric O$_2$, the δ$^{18}$O composition of food and water are highly variable, and thus likely to explain any differences found in the δ$^{18}$O ratios of sympatric taxa. The δ$^{18}$O in plants is more positive than in their source water, which is ultimately derived from local rain. In most cases, liquid water in plant roots and stems is isotopically similar to drinking water available for herbivores, but leaf water is relatively enriched in H$_2^{18}$O due to preferential evapotranspiration of the lighter H$_2^{16}$O molecule. A study of the oxygen isotope composition of modern South African ungulates (Sponheimer and Lee-Thorp, 2001) revealed that mixed-feeding impalas (Aepyceros melampus), grazing tsessebes (Damaliscus lunatus) and blue wildebeests (Connochaetes taurinus) obtain relatively more of their water from green vegetation and are significantly enriched in $^{18}$O compared to other herbivores such as warthog (Phacochoerus aethiopicus), waterbuck (Kobus ellipsiprymnus) and giraffe (Giraffa camelopardalis), which derive more of their total water intake from drinking. Similar differences were detected by Harris and Cerling (2002) between grazing and browsing East African ungulates. Thus, among extinct ungulates a more positive result would indicate that the species obtained most of its water requirements from the plants eaten rather than from drinking. Animal tissues consist mainly of proteins whereas plant tissues consist mainly of carbohydrates. Given that proteins are depleted in $^{18}$O compared to carbohydrates, carnivores show lower δ$^{18}$O values than herbivores (Sponheimer and Lee-Thorp, 2001).
4. Trophic level and palaeodietary inferences

Figure 2 shows a plot of $\delta^{13}$C and $\delta^{15}$N values measured in large mammals from Venta Micena. The range of $\delta^{13}$C values for ungulates (−27‰ to −20‰) agrees with that of modern herbivores eating C₃ plants, which confirms that C₄ grasses were absent from southern Spain during early Pleistocene times (Palmqvist et al., 2003). There are, however, important variations of carbon-isotope ratios among ungulates (Fig. 3), with perissodactyls showing the lowest $\delta^{13}$C values (range: −26.7‰ to −24.2‰) and bovids the highest ones (range: −23.9‰ to −20.1‰). This difference is statistically significant ($t = 19.25, p < 0.0001$) and does not seem to indicate a different feeding behaviour for both groups, given their hypsodonty values (HI). In fact, the two perissodactyl species have similar $\delta^{13}$C values, but the highly hypsodont cheek teeth ($HI = 6.1$) of horse *Equus altidens* identify it as a grazer ($HI = 3.9$-$8.7$ for grazing perissodactyls; Mendoza *et al.*, 2002), whereas the brachydont teeth ($HI = 1.8$) of rhino *Stephanorhinus* sp. indicate that it was a mixed feeder or browser ($HI = 0.8$-$2.2$ for mixed-feeding and browsing perissodactyls; Mendoza *et al.*, 2002). Both species presumably inhabited open, relatively unforested environments, given their comparatively high $\delta^{15}$N values (Fig. 3). Among the bovids, the bovine *Bison* sp. ($HI = 3.9$) and the caprine *Hemitragus albus* ($HI = 4.4$) have moderately hypsodont teeth, indicative of a diet composed mainly of grass ($HI = 3.8$-$6.1$ for modern grazing bovids; Mendoza *et al.*, 2002). In contrast, the ovibovine *Soergelia minor* has mesodont teeth ($HI = 2.9$), suggestive of a mixed diet ($HI = 2.5$-$5.3$ for mixed-feeding bovids from open habitat; Mendoza *et al.*, 2002). The relatively high $\delta^{15}$N values of all these extinct bovids suggest that they dwelled in unforested environments. Cervids *Praemegaceros* cf. *verticornis* and *Pseudodama* sp. also show lower $\delta^{13}$C mean ratios than bovids (−25.9‰ to −22.2‰) ($t = 11.71, p < 0.0001$, two-tailed test). Both species have low-crowned teeth ($HI = 1.6$ and 1.7, respectively), which suggests that they were mixed feeders or browsers in closed habitat, as most cervids ($HI = 1.1$-$2.8$ in modern deer; Mendoza *et al.*, 2002). In fact, *P. verticornis* shows the lowest $\delta^{15}$N contents among ungulates and lower $\delta^{13}$C values than other ruminants (Fig. 3), which suggests a browsing diet in closed canopy.
The similarity in $^{13}\text{C}/^{12}\text{C}$ ratios shown by the two perissodactyls is unlikely to indicate similarity in diets, and shared dietary differences to the ruminants, especially as there were no $\text{C}_4$ grasses in this locality. Rather, it reflects a lower isotope enrichment factor for the heavy-carbon isotope in these monogastric herbivores than in ruminants, related to physiological differences between both groups in their digestive systems (hindgut and foregut, respectively). Of interest for this study, Cerling and Harris (1999) found that Burchell's zebras ($\textit{Equus burchelli}$), whose diet is composed of nearly 100% grass (McNaughton and Georgiadis 1986), are consistently 1-2‰ depleted in $\delta^{13}\text{C}$ values for tooth enamel compared to sympatric ruminant hypergrazers such as alcelaphine bovids; this depletion implies a lower isotope enrichment factor for zebras, resulting from their lower digestive efficiency. A similar difference was detected by Lee-Thorp and Van der Merwe (1987) between bone bioapatite samples from zebra and wildebeest ($\textit{Connochaetes taurinus}$). Recent studies of bone collagen isotopes of European fauna over the last glacial cycle and the Holocene also showed consistently 1-2‰ depleted $\delta^{13}\text{C}$ values in horse compared to ruminants (Bocherens and Drucker, 2003; Richards and Hedges, 2003). However, it is worth noting that the difference in $\delta^{13}\text{C}$ ratios between ruminants and perissodactyls in Venta Micena is greater than the differences reported at other sites, which suggests that some dietary differences must be also involved.

The two megaherbivores, elephant $\textit{Mammuthus meridionalis}$ and hippo $\textit{Hippopotamus antiquus}$, show high $\delta^{13}\text{C}$ values, similar to those of bovids (Fig. 3). The proboscidean seems to have been a mixed feeder like modern African elephants ($\textit{Loxodonta africana}$), although grass probably was a more significant component of its diet according to carbon-isotopes of tooth enamel in fossil $\textit{Mammuthus}$ from Africa and North America (Koch et al., 1998; Cerling et al., 1999). In the case of $H. antiquus$, the modern hippo, $\textit{Hippopotamus amphibious}$, is a reputedly grazer that has brachydont teeth, as in the specimens of Venta Micena. The reason for this apparent anomalous condition of low hypsodonty is most likely related to the fact that hippos have low metabolic rates, consuming less food per day than would be expected for animals of their size (Novak, 1999;
Schwarm et al., 2006), which translates in a lower amount of wear on the teeth (Mendoza et al., 2002). In addition, a recent study of the isotopic composition of enamel in several African populations of *H. amphibious* has shown that modern hippos have a more varied diet than usually thought, including significant amounts of C\textsubscript{3} plants in closed to moderately open environments (Boisserie et al., 2005).

The higher $\delta^{13}$C enrichment of the Venta Micena ruminants, in comparison with the perissodactyls, is probably related to the higher rates of methane production in the forestomach of ruminants than in the hindgut of monogastric herbivores, which derive a lower fraction of maintenance energy from methanogenetic activity of bacteria (Crutzen et al., 1986; Vermorel et al., 1997; Schulze et al., 1998; Metges et al., 1990; Hedges, 2003). However, it seems anomalous that the elephant, a hindgut fermenter, should have an enrichment value that resembles that of the ruminants. However, elephants are not closely related to perissodactyls (Springer et al., 1997) and phylogenetic differences may be at work here. In addition, elephants have a shorter and wider small intestine in comparison to other hindgut fermenters, which is related to the need for animals of such large size to have a very rapid passage rate of the ingesta (Clauss et al., 2003).

The ruminant type of forestomach fermentation provides a clear advantage under conditions of limiting quantities of food: ruminants are very efficient at extracting maximum amounts out of the cellulose and cell contents of food of moderate fibre content, and if feeding on food of relatively good quality they can subsist on a lesser amount of food per day (~70\%) than a hindgut fermenter of similar size, which relies on a rapid passage time and the processing of large quantities of food (Janis, 1976; Janis et al., 1984; Duncan et al., 1990). If food is not a limiting factor, however, hindgut fermentation works well with plants of low nutritive value, such as herbage with high fibre contents. This is because a large volume of food can be processed rapidly and the monogastric herbivore can obtain a large quantity of energy from the cell contents in a short time. Among present-day ungulates, the mid range of body sizes (10–1000 kg) is dominated by ruminants. The exceptions are equids, which can feed on low quality grasses too fibrous for a ruminant to subsist
on, and tapirs, which have remained as a relict group of tropical forest browsers. There are physiological reasons behind the absence of ruminants at very small body sizes: given that the basal metabolic rate scales allometrically to the 0.75 power of body mass (and a similar exponent applies to food intake rate in herbivores), small animals have relatively greater energetic demands than larger ones (Kleiber, 1975; McNab, 1986; Shipley et al., 1994). Ruminants less than 10 kg (tragulids and duikers) eat primarily non-fibrous food items, such as young leaves, buds, seeds and fruit, and the small herbivores that can subsist on more fibrous diets are all hindgut fermenters such as hyraxes and lagomorphs. Physiological scaling effects also operate at the largest body sizes: as the retention time of food in the digestive tract scales to the 0.27 power of body mass (Illius and Gordon, 1992), there is no advantage to foregut fermentation above a certain body mass (600 kg for browse, 1200 kg for grass forage) in terms of digestive efficiency, because the retention times and percentages of fibre digestibility are similar for both foregut and hindgut fermenters (Demment and Van Soest, 1985; Prins and Kreulen, 1991; Justice and Smith, 1992; Van Soest, 1994). In addition, while specific metabolic rate decreases with increasing mass, gut capacity remains a constant fraction of body size (Bell, 1971; Jarman, 1974; Geist, 1974; Parra, 1978; Justice and Smith, 1992). This implies that larger ungulates are able to support their lower specific metabolic requirements by ingesting forage of lower quality (Van Soest, 1996). Ruminants are at a disadvantage at very large body sizes, as they are unable to accelerate the passage rate of their ingesta, and may suffer from other physiological problems such as a decreased capacity for water retention (Clauss et al., 2003).

The observation on the upper size limit for ruminants, based on the range of modern forms, is supported by the fossil record on extinct ruminants and tylopods, which did not, with the possible exception of the sivatheriine giraffids and some Pliocene North American camels, surpass extant species in maximum body size.

The digestive physiology of elephants, however, deviates from the common scheme postulated for herbivores of increasing body mass (Clauss et al., 2003; Loehlein et al., 2003): elephants do not have long ingesta passage rates and achieve only comparatively low digestibility
coefficients. As discussed above, the main nutritional advantage of large body size is that larger animals have lower relative energy requirements and that, due to their increased gastrointestinal tract capacity, they achieve longer ingesta passage rates, which allows them to use forage of lower quality. However, the fermentation of plant material cannot be optimized endlessly, because there is a time when plant fibre is totally fermented and energy losses due to methanogenic bacteria become punitive (Clauss et al., 2003). Therefore, very large herbivores need to evolve adaptations for a comparative acceleration of ingesta passage. Among the extant ungulates, elephants, with their shortened gastrointestinal tract and reduced caecum, are indicators of a trend that allowed even larger hindgut fermenting mammals to exist (Clauss et al., 2003). Foregut fermenting ungulates did not evolve species in which the intake-limiting effect of the foregut could be reduced (e.g., by special bypass structures), and hence their digestive model imposed an intrinsic body size limit for ruminants. This limit will be lower the more the diet enhances the ingesta retention and hence the intake-limiting effect: due to the mechanical characteristics of grass, grazing ruminants cannot become as large as the largest browsing ruminant, the giraffe. In contrast, the design of the gastrointestinal tract of hindgut fermenters allows adaptations for relative passage acceleration, which explains why the largest extinct mammal (Paraceratherium, with a body mass of 10,000-15,000 kg; Fortelius and Kappelman, 1993) was a hindgut fermenter (Clauss et al., 2003).

Figure 3 shows $\delta^{15}$N values for the large mammals from Venta Micena. Carnivores Homotherium latidens, Megantereon whitei, Panthera cf. gombaszoegensis, Pachycrocuta brevirostris, Lycaon lycaonoides, and Canis mosbachensis show higher values than ungulates except in the case of H. antiquus and the sample analyzed of the single specimen of muskoxen, Praeovibos sp., preserved in the assemblage. The isotopic fractionation between carnivores and herbivores is in accordance with the enrichment value expected from increasing one trophic level, indicating that the collagen extracted from the fossils did not undergo diagenetic alteration.

With the only exceptions of hippo and muskoxen, all ungulate species record isotopic values that agree well with a diet of N$_2$-fixing plants (Fig. 3). The high $\delta^{15}$N values obtained for H.
*antiquus*, which are even more elevated than those of the Venta Micena carnivores, suggest that this extinct hippo fed predominantly on aquatic plants, instead of consuming terrestrial grasses as do living hippos (Boisserie et al., 2005). Sealy et al. (1987) found $\delta^{15}$N values in *H. amphibious* similar to those of other sympatric grazing artiodactyls from southern Africa that feed on terrestrial vegetation, and smaller than those in carnivores. The unexpected diet of *H. antiquus* probably relates to the huge size of this extinct hippo: preliminary estimates based on the diaphyseal diameter of major limb bones provide a figure of ~3200 kg of average body mass for *H. antiquus* (1500 kg for *H. amphibius*; Novak, 1999). In addition, *H. antiquus* had shorter metapodials than modern hippos. Given the fact that *H. amphibius* is not well designed for dwelling on land, the enormous size and short limbs of *H. antiquus* must have posed even more severe limitations for terrestrial locomotion.

Modern muskoxen (*Ovibos moschatus*) are sexually dimorphic ruminants that live in a highly seasonal environment, consuming willow, forbs and sedge-dominated vegetation types (Klein and Bay, 1994). However, during winter muskoxen subsist primarily on lichens and some senescent browse. Lichens, although potentially high in digestible energy, contain less protein than required for metabolic maintenance. Thus, the elevated $\delta^{15}$N value of the single specimen of *Praeovibos* sp. (Fig. 3) could indicate increased recycling of nitrogen from body protein during winter due to a poor quality diet (Barboza and Reynolds, 2004; Parker et al., 2005).

Perissodactyls and bovids show more positive $\delta^{15}$N ratios than cervids ($t = 7.36, p < 0.05$) (Fig. 3). This indicates that the cervids would preferably feed in closed habitats, where their low $\delta^{15}$N values would result from soil acidity (Rodière et al., 1996; Gröcke, 1997a). Among the perissodactyls, the comparatively high $\delta^{15}$N values of the browsing rhino, *Stephanorhinus* sp., suggest that this species lived in open habitat, in a similar fashion to the modern black rhino (*Diceros bicornis*). $\delta^{15}$N values for elephant, *M. meridionalis*, are similar to those expected in a large monogastric herbivore (Gröcke and Bocherens, 1996). $\delta^{15}$N values for the horse, *E. altidens*, are also congruent with those of hindgut fermenters of medium size living in an open environment.
Carbon and nitrogen isotopic compositions of collagen provide a proxy to reconstruct ancient trophic webs and especially to decipher the relationships between predators and their potential prey (see review in Bocherens and Drucker, 2003). The wide range of $\delta^{13}C$ and $\delta^{15}N$ values in the Venta Micena carnivores reflects resource partitioning among sympatric predators (Figs. 2-3). For example, sabre-tooth cats *H. latidens* and *M. whitei* have quite distinct isotopic signatures, the former showing the highest $\delta^{13}C$ values among carnivores and the latter the lowest ones (Fig. 3). This suggests that both predators specialized on different types of ungulate prey, which is confirmed by differences in their postcranial anatomy (Anyonge, 1996; Arribas and Palmqvist, 1999; Palmqvist and Arribas, 2001; Palmqvist et al., 2003). The dirk-tooth *M. whitei* had a robust body, a low brachial index (~80%) and short metapodials, features that describe it as an ambush predator of forested habitat. In such an environment, browsing ungulates with depleted $\delta^{13}C$ and $\delta^{15}N$ values would have been the preferred prey. The European jaguar *P. gombaszoegensis* was also an ambusher with a postcranial anatomy similar to that of the extant jaguar (*Panthera onca*). The scimitar-tooth *H. latidens* had comparatively long and slender limbs, and a body size similar to that of a modern lion (*Panthera leo*). The forelimb was more elongated than the hind limb, indicating that the animal probably had a sloping back. This suggests adaptations to carry away large prey.

The claws of *Homotherium* were small, with the exception of a well-developed dewclaw in the first digit of the forefoot. The elongated forelimb (brachial index of ~100%) and small claws suggest increased cursoriality in an open habitat and less prey grappling capability than other sabre-tooth cats (Palmqvist et al., 2003). In such habitat, the prey would be large grazing ruminants and juveniles of megaherbivores, as discussed in depth below.

Canids *L. lycaonoides* and *C. mosbachensis* show intermediate $\delta^{13}C$ values, while the hyena *P. brevirostris*, the bone-collecting agent at this locality, has the highest $\delta^{13}C$ values among carnivores. Results obtained in a comparative ecomorphological study of the craniodental anatomy of modern and Pleistocene canids (Palmqvist et al., 1999, 2002) indicate that the larger canid, *L. lycaonoides*, was a hypercarnivore (>70% vertebrate flesh in diet). The postcranial anatomy of this
species is similar to that of African hunting dogs (*Lycaon pictus*), the only living canids with a tetractyl forelimb, which indicates a coursing behaviour. The medium-sized canid, *C. mosbachensis*, has a craniodental anatomy similar to those of modern coyote (*Canis latrans*), with a talonid basin well-developed in the lower carnassial. This suggests an omnivorous diet. The short-faced hyena, *P. brevirostris*, had a body and skull 20% larger than in modern spotted hyenas (*Crocota crocuta*) and was well-adapted for destroying carcasses and consuming bones (Arribas and Palmqvist, 1998; Palmqvist and Arribas, 2001). This hyaenid differed from other species in the shortening of the distal limb segments, which suggests less coursing abilities, although such shortening would provide greater power and more stability for dismembering and carrying large pieces of ungulate carcasses (Turner and Antón, 1996).

The high $\delta^{18}O$ values of cervid *Pseudodama* sp., bovids *H. albus* and *S. minor*, and bear *U. etruscus* (Fig. 4) suggest that these species obtained most of their water requirements from the vegetation rather than from drinking. In contrast, *M. meridionalis, H. antiquus, Stephanorhinus* sp., *P. verticornis*, and *Praeovibos* sp. exhibit the lowest $\delta^{18}O$ ratios, which indicate greater water dependence for these species. *Bison* sp. and *E. altidens* show intermediate $\delta^{18}O$ values but closer to those of megaherbivores and large deer, which suggest moderate water dependence for both grazing species. These results agree with expectations from their living closest relatives (see review in Palmqvist et al., 2003). For example, goats are well-adapted for arid conditions, obtaining most of their water requirements from the vegetation; such physiological specialization seems to have been developed by the caprine *H. albus* and the ovibovine *S. minor*. Modern fallow deer (*Dama dama*) tolerates more arid conditions than red deer (*Cervus elaphus*), showing a lower water intake rate per kg of body mass, and this was clearly the case for *Pseudodama* sp. The largest monogastric mammals from Africa, black rhino, hippo and elephant, show a greater water dependence than grazing ruminants (Bocherens et al., 1996a), which agrees with the low $^{18}O$ contents of *Stephanorhinus, Hippopotamus* and *Mammuthus*, respectively. In fact, Harris and Cerling (2002) found that tooth enamel samples from hippos and elephants from Queen Elizabeth Park, Uganda,
were consistently depleted in $^{18}$O compared with those from C$_4$-grazing ungulates that obtain most of their water from the plants. All carnivore species in Venta Micena show depleted $\delta^{18}$O values in comparison with the herbivores, as would be predicted from the higher protein content of their diet (Sponheimer and Lee-Thorp, 2001).

5. Predator-prey relationships

The differences in $\delta^{13}$C and $\delta^{15}$N values among carnivores (Figs. 2-3) suggest specific predator-prey relationships in this early Pleistocene community (Palmqvist et al., 1996). For example, *H. latidens* shows the highest $\delta^{13}$C and $\delta^{15}$N values among hypercarnivores, which would indicate that this was the top predator of the palaeocommunity (i.e., the only one able to hunt on very large prey such as juveniles of megafauna and adult ungulates of medium-to-large size). In contrast, *M. whitei* and *P. gombaszoegensis* show the lowest $\delta^{13}$C and $\delta^{15}$N ratios, which may provide evidence that browsing ungulates from forest represented an important fraction of their diet.

Previous studies indicate that primary and secondary productivities of the large mammal fauna from Venta Micena assemblage are balanced, which suggests that all large carnivore species living in the original community were preserved in the bone assemblage collected by the hyenas (Palmqvist et al., 2003). Thus, it seems quite reasonable to assume that the potential prey species for each predator were also preserved in the taphocoenosis. The issue here is to assign the preferred ungulate preys to each predator and to quantify their relative contributions to the predator’s diet.

Given that the only sources of carbon and nitrogen for a carnivore come from its diet, the composition of its tissues will be a function of what the animal ate. Using the isotopic enrichment from prey to predator and the principle of mass balance, the dual linear mixing model (Phillips, 2001) allows estimating quantitatively the proportional contribution of several ungulate prey species to the diet of a carnivore. For two isotopes and three prey sources, their relative abundances in the diet of a predator consuming them may be estimated from the following equations:

$$\delta^{13}C_{\text{predator}} = f_A \delta^{13}C_{\text{prey } A} + f_B \delta^{13}C_{\text{prey } B} + f_C \delta^{13}C_{\text{prey } C},$$
\[
\delta^{15}N_{\text{predator}} = f_A \delta^{15}N'_{\text{prey } A} + f_B \delta^{15}N'_{\text{prey } B} + f_C \delta^{15}N'_{\text{prey } C}, \]

where \(\delta^{13}C'_{\text{prey}}\) and \(\delta^{15}N'_{\text{prey}}\) are the carbon- and nitrogen-isotope ratios of prey after correction for trophic fractionation, and \(f\) represents the relative contributions of preys \(A\), \(B\) and \(C\) to the diet of predator, respectively. This model has been satisfactorily applied to studies of the diet of living carnivores (Phillips, 2001; Phillips and Koch, 2002; Phillips and Gregg, 2003) as well as to derive inferences on the diet of extinct species, including Neanderthals and anatomically modern humans (Drucker and Bocherens, 2004; Bocherens et al., 2005; Phillips et al., 2005). If an apparently unreasonable solution is obtained for the contribution of a given prey species (i.e., \(f < 0\) or \(> 1\)), this can mean either that an important food source was not included in the analysis or that trophic correction factors were not estimated appropriately (Phillips, 2001). However, the model will yield correct results if some of the sources are not in the predator’s diet and even can be used for investigating the dietary composition of a mixture of more than \(n+1\) sources for \(n\) isotopes (Phillips and Gregg, 2003; Bocherens et al., 2005).

This methodological approach was applied to the four hypercarnivore species identified at Venta Micena, \(H. \text{latidens}\), \(M. \text{whitei}\), \(P. \text{gombaszoegensis}\), and \(L. \text{lycaonoides}\). The giant hyena, \(P. \text{brevirostris}\), was excluded from the analysis because taphonomic studies unequivocally indicate that it scavenged the prey of hypercarnivores (Palmqvist et al., 1996; Arribas and Palmqvist, 1998; Palmqvist and Arribas, 2001). The Etruscan bear, \(U. \text{etruscus}\), was not considered because its teeth are similar to those of modern brown bears (\(U. \text{arctos}\)), which is evidence of an omnivorous diet. The craniodental anatomy of the medium-sized canid, \(C. \text{mosbachensis}\), indicates a more omnivorous behaviour than that of the hunting dog \(L. \text{lycaonoides}\) (Palmqvist et al., 1999, 2002) and was also excluded from subsequent analyses. The comparatively low \(\delta^{15}N\) values of this coyote-like species suggest that invertebrates and fruit were also an important fraction of its diet. However, in the case of the bear, its very high \(\delta^{15}N\) values are intriguing. Perhaps this species
consumed regularly fish, in contrast with the other carnivores, or the high δ^{15}N values may have resulted from the physiology of dormancy: Fernández-Mosquera et al. (2001) have reported higher δ^{15}N values in cave bears (*Ursus spelaeus*) that lived during colder periods, which suggest a reuse of urea in synthesizing amino acids with prolonged duration of dormancy.

Among ungulates, *H. antiquus* was discarded as a potential prey given its enormous size and amphibious behaviour, which makes it difficult to conceive of a predator specializing on it; the scarce remains of this species in the assemblage would be the result of the scavenging activity of hyenas on carcasses of adult animals dead by other causes than predation and juveniles hunted occasionally by sabre-tooth cats (Fig. 7B). Elephant *M. meridionalis*, however, was analyzed because isotopic data were available for young individuals, the only age stage that would be susceptible to predation according to data on lion predation on modern elephants (see review in Palmqvist et al., 1996). Finally, *Praeovibos* sp. was discarded because this species, poorly represented in the assemblage, presumably lived in mountainous areas, as modern muskoxen.

The first step is to quantify the trophic fractionation (Δδ^{13}C, Δδ^{15}N) between diet and animal tissues, which not an easy task given that enrichment values depend on the type of food sources and animal tissues analyzed. Trophic fractionation in mink and bear range from −2.2 to 4.9‰ for δ^{13}C, and from 2.3 to 4.1‰ for δ^{15}N, respectively (Phillips and Koch, 2002). In a review of studies developed under experimental conditions on isotopic enrichment between diet and collagen, Bocherens and Drucker (2003) found that the enrichment factor ranges from 3.7‰ to 6.0‰ for carbon and from 1.7‰ to 6.9‰ for nitrogen, respectively. These data agree with the commonly quoted ranges for enrichment values of bone collagen of 0 to 2‰ for carbon and 3 to 5‰ for nitrogen. Figure 5 shows the distribution of differences in δ^{13}C and δ^{15}N mean values for bone collagen between carnivores and their potential ungulate prey in a set (N = 18) of modern ecosystems and fossil assemblages compiled from the literature. Dentine collagen was not considered here because it tends to show higher δ^{15}N values than bone collagen in some species, probably due to a suckling isotopic signal retained in dentine and eliminated in bone (Bocherens et
al., 2001). Although the range of $\Delta^{13}C$ and $\Delta^{15}N$ values in these communities is very wide (Fig. 5), the mean fractionations (1.6‰ and 3.9‰, respectively) are close to those reported previously (Phillips and Koch, 2002; Bocherens and Drucker, 2003).

In the case of Venta Micena, the $\Delta^{13}C$ and $\Delta^{15}N$ values between herbivores and carnivores, weighted according to NISP values, are 0.8‰ and 2.8‰, respectively. Given that both values are close to the modal classes in the reference dataset (Fig. 5), they were used for correcting the $\delta^{13}C$ and $\delta^{15}N$ values of ungulate prey prior to applying the dual linear mixing model to hypercarnivores. It could be argued, however, that it is circular to apply fractionations calculated from the fossil animals whose diets are being reconstructed. However, this is the most reasonable solution, because in this way the centroids for herbivores and carnivores will match in the plot of $\delta^{13}C$ and $\delta^{15}N$ values corrected for trophic fractionation, which guarantees that each predator will be enclosed in a triangle defined by its three main ungulate prey species (Fig. 6). The relative contributions of each ungulate prey to the diet of each hypercarnivore were calculated using the software IsoSource v. 1.3.1 (http://www.epa.gov/wed/pages/models/stableIsotopes/isosource/IsoSourceV1_3_1.zip).

Obviously, there are many possible solutions for the diet of each predator, but the reasonable ones (i.e., those that do not provide negative estimates for one or more prey species) involve only three possible ungulate prey according to the spatial distribution of predators in the $\delta^{13}C - \delta^{15}N$ diagram (see Phillips and Gregg, 2003: Fig. 6). It is worth noting, however, that the software IsoSource allows estimating the dietary contributions of more than three preys. In doing so, the program calculates the mean proportion for each ungulate in the diet of a given predator using the average of the estimates obtained in all the combinations of three potential preys. However, this procedure provides non-unique solutions, which results in increased uncertainty on the source contributions (see discussions in Phillips and Gregg, 2001, 2003; Phillips et al., 2005). In addition, if all ungulate species preserved in the assemblage are considered as a potential prey of each carnivore, this would lead to unrealistic solutions (e.g., to consider that the wild dog L. lycaonoides hunted megafauna or that the large sabre-tooth H. latidens pursued small ungulates; see discussion
in Palmqvist et al., 1996), which would in turn distort the estimates obtained for the more probable preys.

The δ^{13}C - δ^{15}N plot displayed on Figure 6 shows the prey combinations for the four hypercarnivores under the most realistic scenario according to previous ecomorphological studies on the Venta Micena predators and their analogies to modern carnivores (Palmqvist et al., 1996, 2003; Palmqvist and Arribas, 2001). It is worth introducing, however, a cautionary note on the reliability of the contributions of each prey to the diet of each predator, as they would vary if other enrichment factors are used. According to our results (Fig. 6), the preferred preys of H. latidens were grazing and mixed-feeding herbivores from open habitat of medium-to-large size, Bison sp. (52%), E. altidens (38%) and M. meridionalis (10%) (Figs. 7A, C). The likelihood of such specialized hunting behaviour is evident in the case of the related North American species H. serum, known in high numbers from the late Pleistocene site of Friesenhahn cave, a locality interpreted as a sabre-tooth’s den and associated with numerous skeletal remains of adult bison and juvenile mammoths (Marean and Ehrhardt, 1995). The diet of sabre-tooth M. whitei includes E. altidens (59%), P. verticornis (31%) and S. minor (10%) (Figs. 7E, F). The jaguar P. gombaszoegensis preyed upon P. verticornis (43%), Pseudodama sp. (38%) and S. minor (19%).

Finally, the pack hunting L. lycaonoides was likely the most versatile predator of this early Pleistocene community due to its social behaviour (Palmqvist et al., 1999) and had a diet that included E. altidens (58%), H. albus (30%) and Pseudodama sp. (12%), which are low-to-medium sized ungulates from open habitat (Fig. 7D). This combination of prey species is not unreasonable for L. lycaonoides, as modern painted dogs show similar predatory habits in the Serengeti, where Thomson’s gazelle and zebra represent 38% and 20% of prey captures, respectively (Malcom and Van Lawick, 1975).

The distribution of ungulate prey described above, based on isotopic signatures, reflect resource partitioning among sympatric predators in Venta Micena. According with the results obtained, coursing carnivores H. latidens and L. lycaonoides hunted ungulate prey in open habitat,
while *M. whitei* and *P. gombaszoegensis* ambushed their prey in the margins between forest and savannah. This is congruent with the palaeoenvironmental reconstruction of Venta Micena, which is interpreted as a wooded savannah (Mendoza et al., 2005). It is interesting to note that the short-faced hyena, *P. brevirostris*, a species specializing in scavenging the prey of hypercarnivores, shows a $\delta^{15}$N value (Fig. 3) that matches the one expected for a carnivore that consumed all of the ungulates from open habitat preserved in the faunal assemblage (Fig. 7G, H).

Figure 8 shows the average frequencies of ungulate species estimated for a hypothetical death assemblage based on the expectations of the dual linear mixing model. Such frequencies were calculated assuming that: 1) each carnivore exploited the carcasses of its prey to the same degree; and 2) that each predator contributed similar proportions of kills to the death assemblage collected by the hyenas. This theoretical assemblage was compared with the relative frequencies of herbivores in Venta Micena, based on NISP counts after correction for preservational bias related to body size (Arribas and Palmqvist, 1998). In general terms, there are only relatively minor differences between the expected and observed abundance for most ungulate species. This suggests that the hyenas scavenged the ungulate carcasses in the proportions in which they were available and confirms the accuracy of the estimates obtained with the mixing model on the diet of the Venta Micena predators.

6. Conclusions

Patterns of abundance of stable-isotopes of bone collagen ($\delta^{13}$C, $\delta^{15}$N) and bioapatite ($\delta^{18}$O) are a useful proxy for reconstructing the trophic structure of the early Pleistocene large mammal community preserved at Venta Micena, and help also in deciphering the relationships between predators and their potential prey. Carbon-isotope ratios reveal physiologic differences between hindgut and foregut fermenting ungulates related to their digestive systems and the differential assimilation of cellulose, with perissodactyls showing a lower isotopic enrichment than elephants and artiodactyls from open habitat. The low values seen in cervids suggest that they were mixed feeders or browsers in closed habitat. Nitrogen-isotope ratios of carnivore and herbivore species
reflect the isotopic enrichment expected from increasing one trophic level, indicating that the collagen preserved was not diagenetically altered. All ungulate species except the hippo and the muskoxen record isotopic values that agree with a diet of N₂-fixing plants. Cervids show depleted δ¹⁵N values resulting from soil acidity in forest, which confirm a browsing diet in closed habitat. The high nitrogen ratio of H. antiquus suggests that this species fed predominantly on aquatic vegetation. In the case of Praeovibos sp., this mountainous species probably fed on lichens. Oxygen-isotopes of enamel and bone apatite reveal that fallow deer Pseudodama and bovids Hemitragus and Soergelia derived most of their water requirements from the vegetation rather than from drinking. The low δ¹⁸O values of megacerine deer Praemegaceros and of megaherbivores Stephanorhinus, Hippopotamus and Mammutthus suggest a greater degree of water dependence for these species.

Carbon- and nitrogen-isotope ratios in carnivores reflect resource partitioning among sympatric predators, providing interesting clues on predator-prey relationships within this ancient community. The application of the dual linear mixing model allows estimating quantitatively the contribution of several ungulate preys to the diet of each hypercarnivore. Specifically, the scimitar-cat Homotherium, a coursing predator, focused on herbivores from open habitat of relatively large size. The diet of Megantereon and the European jaguar, P. gombaszoegensis, included browsing herbivores from closed habitat such as deer, although horses seem to have been the main prey of the ambushing sabre-tooth. The pack-hunting canid Lycaon consumed grazing ungulates from open habitat such as goat and horse. Finally, carbon- and nitrogen-isotopes of the short-faced hyena Pachycrocuta match the values expected for a carnivore that scavenged all of the ungulates preserved in the faunal assemblage, especially those that lived in open environments.

Acknowledgments

K. Fox-Dobbs, M.J. Kohn, B. Martinez-Navarro, J. McDonald, C. Nedin, C. Trueman, S.L. Wing and an anonymous reviewer provided insightful comments on an earlier version of the manuscript. Funding and analytical facilities for biogeochemical analyses were provided by the
University of Málaga, Royal Holloway University of London and at the Stable-Isotope Biogeochemistry Laboratory at McMaster University.

References


Barboza, P.S., Reynolds, P.R. 2004. Monitoring nutrition of a large grazer: muskoxen on the Arctic


collagen: case studies from recent and ancient terrestrial ecosystems. International Journal of
Osteoarchaeology 13, 46–53.

Bocherens, H., Billiou, D., Patou-Mathis, M., Otte, M., Bonjean, D., Toussaint, M., Mariotti, A.
1999. Palaeoenvironmental and palaeodietary implications of isotopic biogeochemistry of late
interglacial Neanderthal and mammal bones in Scladina Cave (Belgium). Journal of

Bocherens, H., Billiou, D., Mariotti, A., Toussaint, M., Patou-Mathis, M., Bonjean, D., Otte, M.
Evolution 40, 497–505.

from isotopic signatures in Pleistocene cave fauna of Southern England. Journal of

($^{13}$C, $^{18}$O) and mammalian enamel from African Pleistocene hominid sites. Palaios 11, 306–318.

($^{13}$C, $^{15}$N) in collagen and soft tissues from Pleistocene mammals from Yakutia: implications for
the palaeobiology of the mammoth steppe. Palaeogeography, Palaeoclimatology, Palaeoecology
126, 31–44.


Figure 1. A: average values of trace-elements (Fe, Mg, Na, Mn, and Sr, in ppm), and B: bulk rock stable-isotope ratios ($\delta^{13}$C and $\delta^{18}$O, in ‰, $^{87}$Sr/$^{86}$Sr) in the samples of sediment collected from the stratigraphic column of the early Pleistocene locality of Venta Micena. C: reconstruction of paleoenvironmental changes in the lake at Venta Micena during early Pleistocene times, inferred from the abundance of trace elements and bulk-rock $\delta^{18}$O analyses in samples collected through the Venta Micena stratigraphic section.

Figure 2. Plot of $\delta^{13}$C and $\delta^{15}$N values of collagen material extracted from bone samples of large mammal species preserved in the lower Pleistocene site of Venta Micena (data from Table 2). The lines represent one standard deviation around the mean for those species in which at least two measurements of stable-isotopes were available.

Figure 3. Box diagrams of $\delta^{13}$C and $\delta^{15}$N values of collagen material extracted from bone samples of large mammal species preserved in Venta Micena (data from Table 2).

Figure 4. Box diagrams of $\delta^{18}$O values of hydroxylapatite from bone samples and tooth enamel of large mammal species preserved in Venta Micena (data from Table 2).

Figure 5. Histograms showing the distribution of mean values of isotopic enrichment of carbon- and nitrogen-isotopes ($\Delta\delta^{13}$C and $\Delta\delta^{15}$N, respectively) in bone collagen between mammalian carnivores and their potential prey for several living communities and fossil assemblages. Arrows indicate the values for Venta Micena (VM), obtained averaging the species means according to NISP values ($H. antiquus$ and $Praeovibos$ sp. excluded). Fossil assemblages (late Pleistocene to Holocene): Bocherens et al., 1995 (Kent’s Cave, England); Fizet et al., 1995.
(Marillac, France); Bocherens et al., 1996 (Yakutia, Russia); Gröcke 1997a (Henschke Cave, Australia); Bocherens et al., 1999 (Scladina Cave, Level 4, Belgium); Bocherens et al., 2001 (Scladina Cave, Levels 1A-1B, Belgium); McNulty et al., 2002 (Natural Trap Cave, USA); Bocherens and Drucker, 2003 (Saint Germain – la Rivière, France; Les Jamblancs, France); Coltrain et al., 2004 (Rancho La Brea, USA); Drucker and Bocherens, 2004 (Saint Césaire – Camiac – La Berbie, France). Recent communities: Ambrose and DeNiro, 1986 (East Africa); Sealy et al., 1987 (Kasungu National Park, Malawi); Van der Merwe, 1989 (South Africa); Schwarcz, 1991 (Ontario, Canada); Sillen and Lee-Thorp, 1994 (Southwestern Cape, South Africa); Bocherens et al., 1996b (Yakutia, Russia); Szepanski et al., 1999 (Alaska, USA); Bocherens and Drucker, 2003 (Bielowiecza forest, Poland).

**Figure 6.** Plot of $\delta^{13}$C and $\delta^{15}$N values for carnivore and ungulate species from Venta Micena, corrected for trophic fractionation ($\Delta \delta^{13}$C = 0.8‰, $\Delta \delta^{15}$N = 2.8‰). Each of the four hypercarnivore species lies within the triangle defined by its three most probable prey species. According to the linear mixing model (Phillips, 2001), the contribution to carnivore diet of each prey defining a vertex in the triangle is obtained as the distance from this vertex to the opposed side in relation to its distance to the predator (both measured on the line connecting prey and predator). Black circles: ungulates; white circles: hypercarnivores; gray circles: omnivores and bone-cracking hyena.

**Figure 7.** Reconstruction of the predatory behaviour of the early Pleistocene hypercarnivores preserved in Venta Micena, according to biogeochemical and ecomorphological inferences. Machairodont *Homotherium latidens*, a pursuit predator, hunted subadult elephants (A) and juveniles of other megaherbivore species such as hippo (B), but these prey represented a minor fraction of its diet, which predominantly included ungulates of medium-to-large size from open environment such as bison (C) and horse. Saber-tooth *Megantereon whitei*, an ambusher, hunted
ungulates from open habitat such as horse (D) and browsing deer from forest (E). The hypercarnivorous canid *Lycaon lycaonoides* probably had a pack-hunting behaviour similar to that of modern African painted dogs, with small-to-medium sized ungulates such as goat as their main prey (F). The giant hyena, *Pachycrocuta brevirostris*, scavenged ungulate carcasses left by the hypercarnivores (G) and transported skeletal elements to the denning site (H), a conclusion also supported by taphonomic evidence (Palmqvist et al., 1996; Arribas and Palmqvist, 1998; Palmqvist and Arribas, 2001). Drawings by Mauricio Antón.

**Figure 8.** Comparison between the relative abundance of ungulates in the large mammals assemblage from Venta Micena, corrected for taphonomic bias resulting from differences in body mass (Arribas and Palmqvist, 1998), and the relative frequencies in which these species were hunted by the four hypercarnivores, deduced in this article from the application of the linear mixing model and assuming that each predator contributed similarly to the kill assemblage collected by the hyenas.
Table 1. Height (m), trace-element abundance (ppm) and bulk-rock stable-isotope ratios ($\delta^{13}$C, $\delta^{18}$O, in $\%$; $^{87}$Sr/$^{86}$Sr) of sedimentary samples (vm-1 to vm-10) collected through the stratigraphic section of Venta Micena.

<table>
<thead>
<tr>
<th>sample</th>
<th>height</th>
<th>Fe</th>
<th>Mg</th>
<th>Na</th>
<th>Mn</th>
<th>Sr</th>
<th>Ca</th>
<th>$\delta^{13}$C</th>
<th>$\delta^{18}$O</th>
<th>$^{87}$Sr/$^{86}$Sr</th>
</tr>
</thead>
<tbody>
<tr>
<td>vm-11</td>
<td>3.85</td>
<td>897</td>
<td>3975</td>
<td>500</td>
<td>40</td>
<td>979</td>
<td>303917</td>
<td>-7.36</td>
<td>-5.57</td>
<td>0.707836</td>
</tr>
<tr>
<td>vm-10</td>
<td>3.35</td>
<td>2143</td>
<td>5556</td>
<td>682</td>
<td>22</td>
<td>1197</td>
<td>279964</td>
<td>-6.98</td>
<td>-5.40</td>
<td>0.707857</td>
</tr>
<tr>
<td>vm-9</td>
<td>3.15</td>
<td>1007</td>
<td>6068</td>
<td>3245</td>
<td>40</td>
<td>831</td>
<td>297197</td>
<td>-7.42</td>
<td>-5.75</td>
<td>0.707826</td>
</tr>
<tr>
<td>vm-8</td>
<td>2.50</td>
<td>2486</td>
<td>6202</td>
<td>334</td>
<td>50</td>
<td>327</td>
<td>195090</td>
<td>-5.19</td>
<td>-5.18</td>
<td>0.708020</td>
</tr>
<tr>
<td>vm-7</td>
<td>1.95</td>
<td>1695</td>
<td>7396</td>
<td>438</td>
<td>107</td>
<td>435</td>
<td>256736</td>
<td>-5.60</td>
<td>-4.80</td>
<td>0.707972</td>
</tr>
<tr>
<td>vm-6</td>
<td>1.65</td>
<td>2526</td>
<td>8101</td>
<td>389</td>
<td>105</td>
<td>414</td>
<td>236984</td>
<td>-5.23</td>
<td>-5.18</td>
<td>0.707990</td>
</tr>
<tr>
<td>vm-5</td>
<td>1.40</td>
<td>4551</td>
<td>8146</td>
<td>554</td>
<td>120</td>
<td>447</td>
<td>224936</td>
<td>-5.31</td>
<td>-4.93</td>
<td>0.707989</td>
</tr>
<tr>
<td>vm-4</td>
<td>1.15</td>
<td>2256</td>
<td>6465</td>
<td>430</td>
<td>93</td>
<td>449</td>
<td>260913</td>
<td>-5.61</td>
<td>-5.37</td>
<td>0.707947</td>
</tr>
<tr>
<td>vm-3</td>
<td>1.05</td>
<td>5366</td>
<td>8768</td>
<td>754</td>
<td>95</td>
<td>555</td>
<td>202355</td>
<td>-5.53</td>
<td>-4.31</td>
<td>0.708008</td>
</tr>
<tr>
<td>vm-2</td>
<td>0.85</td>
<td>4401</td>
<td>17445</td>
<td>1019</td>
<td>48</td>
<td>647</td>
<td>137997</td>
<td>-6.34</td>
<td>-4.59</td>
<td>0.707891</td>
</tr>
<tr>
<td>vm-1</td>
<td>0.40</td>
<td>4400</td>
<td>68451</td>
<td>1656</td>
<td>136</td>
<td>313</td>
<td>126781</td>
<td>-4.12</td>
<td>1.79</td>
<td>0.707949</td>
</tr>
</tbody>
</table>
Table 2. C:N proportions and isotopic ratios from collagen ($\delta^{13}$C, $\delta^{15}$N) and hydroxylapatite ($\delta^{18}$O) extracted from bone and tooth samples of large mammal species preserved in the early Pleistocene locality of Venta Micena (juv.: juvenile, subad.: subadult; ad.: adult).

<table>
<thead>
<tr>
<th>Species</th>
<th>Sample code</th>
<th>Fossil specimen (tooth or bone portion)</th>
<th>$\delta^{13}$C$_{\text{col.}}$</th>
<th>$\delta^{15}$N$_{\text{stat.}}$</th>
<th>$\delta^{18}$O$_{\text{stat.}}$</th>
<th>C:N$_{\text{stat.}}$</th>
</tr>
</thead>
<tbody>
<tr>
<td>Mammuthus meridionalis (juv.)</td>
<td>VM-4439</td>
<td>proximal metacarpal</td>
<td>-21.4</td>
<td>+4.7</td>
<td>-4.5</td>
<td>3.1</td>
</tr>
<tr>
<td>Mammuthus meridionalis (juv.)</td>
<td>VM-OX1</td>
<td>enamel fragment of deciduous tooth</td>
<td>--</td>
<td>--</td>
<td>-3.9</td>
<td>--</td>
</tr>
<tr>
<td>Mammuthus meridionalis (juv./subad.)</td>
<td>VM-3581</td>
<td>vertebrae fragment (neural arch)</td>
<td>-20.6</td>
<td>+2.9</td>
<td>-2.8</td>
<td>3.3</td>
</tr>
<tr>
<td>Mammuthus meridionalis (juv./subad.)</td>
<td>VM-3581b</td>
<td>vertebrae fragment (neural arch)</td>
<td>-23.4</td>
<td>+4.3</td>
<td>-5.3</td>
<td>3.6</td>
</tr>
<tr>
<td>Mammuthus meridionalis (subad./ad.)</td>
<td>VM-2172</td>
<td>carpal bone</td>
<td>-21.1</td>
<td>+3.2</td>
<td>-5.9</td>
<td>3.4</td>
</tr>
<tr>
<td>Mammuthus meridionalis (subad./ad.)</td>
<td>VM-4482</td>
<td>carpal bone</td>
<td>--</td>
<td>--</td>
<td>-4.7</td>
<td>--</td>
</tr>
<tr>
<td>Mammuthus meridionalis (juv.)</td>
<td>VM-4607</td>
<td>cranial fragment of a newborn individual or fetus</td>
<td>-22.8</td>
<td>+3.5</td>
<td>-4.5</td>
<td>3.5</td>
</tr>
<tr>
<td>Mammuthus meridionalis (juv.)</td>
<td>VM-4607b</td>
<td>cranial fragment of a newborn individual or fetus</td>
<td>-22.9</td>
<td>+3.2</td>
<td>-4.5</td>
<td>3.2</td>
</tr>
<tr>
<td>Mammuthus meridionalis (juv.)</td>
<td>VM-84-C3-C8-49</td>
<td>left maxilla with dp$^3$ and dp$^4$ partially worn</td>
<td>--</td>
<td>--</td>
<td>-4.5</td>
<td>--</td>
</tr>
<tr>
<td>Mammuthus meridionalis (juv.)</td>
<td>VM-4439b</td>
<td>proximal metacarpal</td>
<td>-23.4</td>
<td>+3.8</td>
<td>-5.3</td>
<td>3.2</td>
</tr>
<tr>
<td>Mammuthus meridionalis (juv.)</td>
<td>VM-3674</td>
<td>vertebrae fragment</td>
<td>--</td>
<td>--</td>
<td>-5.0</td>
<td>--</td>
</tr>
<tr>
<td>cf. Mammuthus meridionalis (juv.)</td>
<td>VM-4332</td>
<td>right humeral diaphysis of a newborn individual</td>
<td>--</td>
<td>--</td>
<td>-1.4</td>
<td>--</td>
</tr>
<tr>
<td>cf. Mammuthus meridionalis (juv.)</td>
<td>VM-1919</td>
<td>left humeral diaphysis of a newborn individual</td>
<td>-22.2</td>
<td>+5.1</td>
<td>-5.6</td>
<td>3.6</td>
</tr>
</tbody>
</table>

Mean for Mammuthus (N = 8/13)

Hippopotamus antiquus (ad.)

Hippopotamus antiquus (ad.)

Hippopotamus antiquus (ad.)

Hippopotamus antiquus (ad.)

Hippopotamus antiquus (ad.)

Hippopotamus antiquus (ad.)

Hippopotamus antiquus (juv.)

Hippopotamus antiquus (juv.)

Hippopotamus antiquus (juv.)

Hippopotamus antiquus (juv.)

Hippopotamus antiquus (juv.)

Hippopotamus antiquus (juv.)

Mean for Hippopotamus (N = 5/7)

Bison sp. (ad.)

Bison sp. (ad.)

Bison sp. (ad.)

Bison sp. (ad.)

Bison sp. (ad.)

Bison sp. (ad.)

Bison sp. (ad.)

Bison sp. (ad.)

Bison sp. (ad.)

Bison sp. (ad.)

Bison sp. (juv.)

Bison sp. (juv./subad.)

Mean for Bison (N = 6/11)

Praeovibos sp. (ad.)

Soergelia minor (ad.)

$\delta^{13}$C$_{\text{col.}}$ = -22.26 ± 0.67, $\delta^{15}$N$_{\text{stat.}}$ = -6.78 ± 0.88, $\delta^{18}$O$_{\text{stat.}}$ = -5.50 ± 1.96, C:N$_{\text{stat.}}$ = 3.40 ± 0.11

Mean for Praeovibos (N = 2/2)

Praeovibos sp. (ad.)

Soergelia minor (ad.)

$\delta^{13}$C$_{\text{col.}}$ = -21.98 ± 0.51, $\delta^{15}$N$_{\text{stat.}}$ = +3.77 ± 0.51, $\delta^{18}$O$_{\text{stat.}}$ = -3.85 ± 1.08, C:N$_{\text{stat.}}$ = 3.37 ± 0.11
<table>
<thead>
<tr>
<th>Specimen No.</th>
<th>Species</th>
<th>Description</th>
<th>CR1</th>
<th>CR2</th>
<th>CR3</th>
<th>CR4</th>
</tr>
</thead>
<tbody>
<tr>
<td>VM-3867</td>
<td>Soergelia minor (ad.)</td>
<td>left metatarsal</td>
<td>-23.2</td>
<td>+3.2</td>
<td>-2.1</td>
<td>3.4</td>
</tr>
<tr>
<td>VM-3867b</td>
<td>Soergelia minor (ad.)</td>
<td>left metatarsal</td>
<td>-23.5</td>
<td>+3.4</td>
<td>-2.3</td>
<td>3.5</td>
</tr>
<tr>
<td>VM-OX10</td>
<td>Soergelia minor (ad.)</td>
<td>enamel fragment of permanent tooth</td>
<td>--</td>
<td>--</td>
<td>-2.6</td>
<td>--</td>
</tr>
<tr>
<td>VM-3982</td>
<td>Soergelia minor (ad.)</td>
<td>left proximal metatarsal</td>
<td>-23.0</td>
<td>+3.8</td>
<td>-2.7</td>
<td>3.3</td>
</tr>
<tr>
<td>VM-4597</td>
<td>Soergelia minor (subad.)</td>
<td>maxilla with unworn P1-P4 and M1, M2, M3 erupting</td>
<td>-23.9</td>
<td>+2.2</td>
<td>--</td>
<td>3.6</td>
</tr>
<tr>
<td>VM-4336</td>
<td>Soergelia minor (ad.)</td>
<td>right horn base</td>
<td>-23.4</td>
<td>+4.3</td>
<td>-2.3</td>
<td>3.6</td>
</tr>
<tr>
<td>VM-3802</td>
<td>Hemitragus albus (ad.)</td>
<td>left distal metatarsal</td>
<td>-20.9</td>
<td>+4.0</td>
<td>-3.4</td>
<td>3.3</td>
</tr>
<tr>
<td>VM-3922</td>
<td>Hemitragus albus (ad.)</td>
<td>right distal metacarpal</td>
<td>-20.1</td>
<td>+3.9</td>
<td>-1.5</td>
<td>3.4</td>
</tr>
<tr>
<td>VM-OX11</td>
<td>Hemitragus albus (ad.)</td>
<td>enamel fragment of permanent tooth</td>
<td>--</td>
<td>--</td>
<td>-3.3</td>
<td>--</td>
</tr>
<tr>
<td>VM-3157</td>
<td>Hemitragus albus (ad.)</td>
<td>distal metatarsal</td>
<td>--</td>
<td>--</td>
<td>-2.1</td>
<td>--</td>
</tr>
<tr>
<td>VM-3449</td>
<td>Hemitragus albus (ad.)</td>
<td>left distal humerus</td>
<td>-20.4</td>
<td>+3.7</td>
<td>-2.9</td>
<td>3.5</td>
</tr>
<tr>
<td>VM-4541a</td>
<td>Hemitragus albus (ad.)</td>
<td>distal right humerus with fused epiphysis</td>
<td>-21.6</td>
<td>+3.4</td>
<td>-3.6</td>
<td>3.6</td>
</tr>
<tr>
<td>VM-3055</td>
<td>Pseudodama sp. (ad.)</td>
<td>right proximal metatarsal</td>
<td>--</td>
<td>--</td>
<td>-3.2</td>
<td>--</td>
</tr>
<tr>
<td>VM-3482</td>
<td>Pseudodama sp. (ad.)</td>
<td>left proximal metacarpal</td>
<td>-23.6</td>
<td>+2.5</td>
<td>-2.6</td>
<td>3.5</td>
</tr>
<tr>
<td>VM-4330</td>
<td>Pseudodama sp. (ad.)</td>
<td>left proximal metacarpal</td>
<td>--</td>
<td>--</td>
<td>-2.4</td>
<td>--</td>
</tr>
<tr>
<td>VM-OX9</td>
<td>Pseudodama sp. (ad.)</td>
<td>enamel fragment of permanent tooth</td>
<td>--</td>
<td>--</td>
<td>-2.6</td>
<td>--</td>
</tr>
<tr>
<td>VM-4410</td>
<td>Pseudodama sp. (ad.)</td>
<td>fragment of left hemimandible with M4-M5</td>
<td>--</td>
<td>--</td>
<td>-2.9</td>
<td>--</td>
</tr>
<tr>
<td>VM-3047</td>
<td>Pseudodama sp. (ad.)</td>
<td>right distal metacarpal</td>
<td>-23.8</td>
<td>+2.9</td>
<td>-3.0</td>
<td>3.3</td>
</tr>
<tr>
<td>VM-3060</td>
<td>Pseudodama sp. (ad.)</td>
<td>right proximal metacarpal</td>
<td>-23.2</td>
<td>+2.6</td>
<td>-2.5</td>
<td>3.4</td>
</tr>
<tr>
<td>VM-4409</td>
<td>Pseudodama sp. (subad.)</td>
<td>left mandible with dp1, heavily worn and M1</td>
<td>-22.8</td>
<td>+1.8</td>
<td>-2.8</td>
<td>3.2</td>
</tr>
<tr>
<td>VM-4037</td>
<td>Pseudodama sp. (subad.)</td>
<td>right mandible with dp1-dp4 and M1 slightly worn</td>
<td>-22.2</td>
<td>+2.2</td>
<td>-1.3</td>
<td>3.0</td>
</tr>
<tr>
<td>VM-3297</td>
<td>Praemegaceros verticornis (ad.)</td>
<td>fragment of right metacarpal diaphysis</td>
<td>-25.9</td>
<td>+1.6</td>
<td>-3.3</td>
<td>3.2</td>
</tr>
<tr>
<td>VM-4155</td>
<td>Praemegaceros verticornis (ad.)</td>
<td>left proximal radius</td>
<td>--</td>
<td>--</td>
<td>-2.8</td>
<td>--</td>
</tr>
<tr>
<td>VM-3111</td>
<td>Praemegaceros verticornis (ad.)</td>
<td>distal metacarpal</td>
<td>--</td>
<td>--</td>
<td>-4.0</td>
<td>--</td>
</tr>
<tr>
<td>VM-3556</td>
<td>Praemegaceros verticornis (ad.)</td>
<td>right distal humerus</td>
<td>-25.6</td>
<td>+1.3</td>
<td>-4.2</td>
<td>3.6</td>
</tr>
<tr>
<td>VM-OX8</td>
<td>Praemegaceros verticornis (ad.)</td>
<td>enamel fragment of permanent tooth</td>
<td>--</td>
<td>--</td>
<td>-4.6</td>
<td>--</td>
</tr>
<tr>
<td>VM-4181</td>
<td>Praemegaceros verticornis (ad.)</td>
<td>left distal humerus</td>
<td>-25.4</td>
<td>+1.8</td>
<td>-3.7</td>
<td>3.4</td>
</tr>
<tr>
<td>VM-3111b</td>
<td>Praemegaceros verticornis (ad.)</td>
<td>distal metacarpal</td>
<td>--</td>
<td>--</td>
<td>-3.9</td>
<td>--</td>
</tr>
<tr>
<td>VM-3224</td>
<td>Praemegaceros verticornis (ad.)</td>
<td>left distal metatarsal</td>
<td>-25.8</td>
<td>+1.4</td>
<td>-3.1</td>
<td>3.5</td>
</tr>
<tr>
<td>VM-3780</td>
<td>Praemegaceros verticornis (ad.)</td>
<td>left distal humerus</td>
<td>--</td>
<td>--</td>
<td>-4.0</td>
<td>--</td>
</tr>
<tr>
<td>VM-84C3-E10-63</td>
<td>Praemegaceros verticornis (juv.)</td>
<td>left mandible with dp2-dp4</td>
<td>-25.9</td>
<td>+2.3</td>
<td>-3.5</td>
<td>3.4</td>
</tr>
<tr>
<td>VM-82-9</td>
<td>Praemegaceros verticornis (juv.)</td>
<td>left maxilla with dp1, dp2 and dp4</td>
<td>-24.4</td>
<td>+1.7</td>
<td>--</td>
<td>3.3</td>
</tr>
<tr>
<td>VM-4039</td>
<td>Praemegaceros verticornis (juv.)</td>
<td>right mandible with dp4 and M1 erupting</td>
<td>-25.2</td>
<td>+2.0</td>
<td>-3.7</td>
<td>3.6</td>
</tr>
<tr>
<td>VM-4394</td>
<td>Praemegaceros verticornis (juv.)</td>
<td>fragment of right maxilla, with dp2, dp3, dp4 and M1</td>
<td>--</td>
<td>--</td>
<td>-2.0</td>
<td>--</td>
</tr>
<tr>
<td>VM-4487</td>
<td>Stephanorhinus etruscus (ad.)</td>
<td>left proximal ulna</td>
<td>-26.5</td>
<td>+3.9</td>
<td>-3.7</td>
<td>3.5</td>
</tr>
<tr>
<td>VM-4510</td>
<td>Stephanorhinus etruscus (ad.)</td>
<td>right humeral diaphysis</td>
<td>-26.6</td>
<td>+3.7</td>
<td>-3.7</td>
<td>3.3</td>
</tr>
<tr>
<td>VM-OX4</td>
<td>Stephanorhinus etruscus (ad.)</td>
<td>tooth fragment</td>
<td>--</td>
<td>--</td>
<td>-3.7</td>
<td>--</td>
</tr>
<tr>
<td>VM-OX5</td>
<td>Stephanorhinus etruscus (ad.)</td>
<td>enamel fragment of permanent tooth</td>
<td>--</td>
<td>--</td>
<td>-5.1</td>
<td>--</td>
</tr>
<tr>
<td>VM-3616</td>
<td>Stephanorhinus etruscus (ad.)</td>
<td>right distal femur</td>
<td>-26.2</td>
<td>+3.5</td>
<td>-4.2</td>
<td>3.3</td>
</tr>
<tr>
<td>VM-3578</td>
<td>Stephanorhinus etruscus (ad.)</td>
<td>left scapula</td>
<td>--</td>
<td>--</td>
<td>-4.6</td>
<td>--</td>
</tr>
<tr>
<td>VM-3578b</td>
<td>Stephanorhinus etruscus (ad.)</td>
<td>left scapula</td>
<td>--</td>
<td>--</td>
<td>-4.8</td>
<td>--</td>
</tr>
<tr>
<td>VM-3578c</td>
<td>Stephanorhinus etruscus (ad.)</td>
<td>left scapula</td>
<td>-25.8</td>
<td>+3.0</td>
<td>-5.2</td>
<td>3.2</td>
</tr>
<tr>
<td>VM-3744</td>
<td>Stephanorhinus etruscus (ad.)</td>
<td>right femoral diaphysis</td>
<td>-24.7</td>
<td>--</td>
<td>-4.3</td>
<td>3.1</td>
</tr>
<tr>
<td>VM-1908</td>
<td>Stephanorhinus etruscus (juv./subad.)</td>
<td>right distal radius with the epiphysis unfused</td>
<td>-25.2</td>
<td>+3.6</td>
<td>-4.8</td>
<td>3.6</td>
</tr>
</tbody>
</table>
### Stephanorhinus etruscus (juv.)
- VM-non coded: fragment of left maxilla, with dP² and dP³

Mean for Stephanorhinus (N = 6/11):

<table>
<thead>
<tr>
<th>Specimen</th>
<th>Description</th>
<th>Mass</th>
<th>Width</th>
<th>Height</th>
<th>p-value 1</th>
<th>p-value 2</th>
<th>p-value 3</th>
<th>p-value 4</th>
</tr>
</thead>
<tbody>
<tr>
<td>VM-3028</td>
<td>right metatarsal lacking the distal epiphysis</td>
<td>-26.1</td>
<td>±2.0</td>
<td>-2.5</td>
<td>3.2</td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>VM-3119</td>
<td>right scapula</td>
<td>-26.7</td>
<td>±3.5</td>
<td>-2.8</td>
<td>3.6</td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>VM-3162</td>
<td>right distal tibia</td>
<td>-26.0</td>
<td>±3.0</td>
<td>-3.0</td>
<td>3.5</td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>VM-3258</td>
<td>right distal tibia</td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>VM-3430</td>
<td>left distal metatarsal</td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>VM-3529</td>
<td>left proximal scapula</td>
<td>-26.0</td>
<td>±3.3</td>
<td>-3.3</td>
<td>3.4</td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>VM-4189</td>
<td>fragment of metatarsal diaphysis</td>
<td>-25.5</td>
<td>±3.5</td>
<td>-2.3</td>
<td>3.5</td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>VM-4421</td>
<td>left proximal radius</td>
<td>-25.0</td>
<td>±2.1</td>
<td>-3.8</td>
<td>3.4</td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>VM-OX2</td>
<td>enamel fragment of permanent tooth</td>
<td></td>
<td></td>
<td></td>
<td></td>
<td>-3.4</td>
<td></td>
<td></td>
</tr>
<tr>
<td>VM-OX3</td>
<td>enamel fragment of permanent tooth</td>
<td></td>
<td></td>
<td></td>
<td></td>
<td>-5.2</td>
<td></td>
<td></td>
</tr>
<tr>
<td>VM-3089</td>
<td>left distal metacarpal</td>
<td></td>
<td></td>
<td></td>
<td></td>
<td>-3.5</td>
<td></td>
<td></td>
</tr>
<tr>
<td>VM-4421b</td>
<td>left proximal radius</td>
<td>-25.8</td>
<td>±2.8</td>
<td>-3.9</td>
<td>3.4</td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>VM-3279</td>
<td>left tibia diaphysis</td>
<td>-26.6</td>
<td>±3.1</td>
<td>-4.3</td>
<td>3.5</td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>VM-3428</td>
<td>distal diaphysis of metatarsal</td>
<td></td>
<td></td>
<td></td>
<td></td>
<td>-2.8</td>
<td></td>
<td></td>
</tr>
<tr>
<td>VM-4403</td>
<td>maxilla with worn dP²-dP³, M²-M' erupting</td>
<td>-25.8</td>
<td>±3.5</td>
<td>-2.5</td>
<td>3.6</td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>VM-602</td>
<td>metacarpal with the proximal epiphysis unfused</td>
<td>-26.3</td>
<td>±2.0</td>
<td>-3.7</td>
<td>3.1</td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>VM-2073</td>
<td>right metacarpal, two thirds of proximal diaphysis</td>
<td></td>
<td></td>
<td></td>
<td></td>
<td>-4.3</td>
<td></td>
<td></td>
</tr>
<tr>
<td>VM-3388</td>
<td>right mandible fragment, with unworn dP₂-dP₄</td>
<td>-25.1</td>
<td>±2.6</td>
<td>-5.2</td>
<td>3.5</td>
<td></td>
<td></td>
<td></td>
</tr>
</tbody>
</table>

Mean for Equus (N = 11/18):

Mean for Pachycrocuta (N = 4/6):

Mean for Homotherium (N = 3/3):

Mean for Megantereon whitei (N = 2/2):

### Meganteine robertsae (ad.)
- VM-84-C3-H5-6: left mandible with dP₂ and dP₃

### Ursus etruscus (ad.)
- VM-1172: right distal radius
- VM-2972: right tibia of an ad. individual
- VM-1903: left radius of an ad. individual
- VM-non coded: skull

Mean for Ursus (N = 3/4):

Mean for Lycaon (N = 3/3):

Mean for Canis (N = 3/3):